



From ocean to coast: Past and future marine climate changes off southeast Australia

Eric C. J. Oliver^{1,2}, Simon J. Wotherspoon¹, Matthew A. Chamberlain³, and Neil J. Holbrook^{1,2}

¹ Institute for Marine and Antarctic Studies, University of Tasmania, Hobart, AUS
² Australian Research Council Centre of Excellence for Climate System Science
² CSIRO Marine and Atmospheric Research (CMAR), Hobart, AUS

translating **nature** into **knowledge**



From Ocean to Coast



Global



Southeastern Australia



• Regional oceanography of the Tasman Sea

RCTIC STUDIES

• western boundary current, eddy-rich region, complex bathymetry







- Global marine climate is warming
- The SW Pacific (Tasman Sea) is a hotspot of change
- Impacts on marine ecology are already being felt
- Historical marine climate in Eastern Tasmania is poorly understood
- Large-scale models and reanalysis are <u>not designed for near-shore</u> <u>studies</u>
- Alternatives
 - Statistical downscaling
 - Dynamical downscaling







- Global marine climate is warming
- The SW Pacific (Tasman Sea) is a hotspot of change
- Impacts on marine ecology are already being felt
- Historical marine climate in Eastern Tasmania is poorly understood
- Large-scale models and reanalysis are <u>not designed for near-shore</u> <u>studies</u>
- Alternatives
 - Statistical downscaling
 - Dynamical downscaling







38

40

42

44



- Global marine climate is warming
- The SW Pacific (Tasman Sea) is a hotspot of change
- Impacts on marine ecology are already being felt
- Historical marine climate in Eastern Tasmania is poorly understood
- Large-scale models and reanalysis are <u>not designed for near-shore</u> <u>studies</u>
- Alternatives
 - Statistical downscaling
 - Dynamical downscaling







- Global marine climate is warming
- The SW Pacific (Tasman Sea) is a hotspot of change
- Impacts on marine ecology are already being felt
- Historical marine climate in Eastern Tasmania is poorly understood
- Large-scale models and reanalysis are <u>not designed for near-shore</u> <u>studies</u>
- Alternatives
 - Statistical downscaling
 - Dynamical downscaling



Change in visible surface kelp canopy (Macrocystis pyrifera)

Longitude 148





Ocean Modeling



- Eddy-resolving dynamical downscaling in Australia region performed by Chamberlain et al. (2010):
- Two <u>ocean model runs</u> using Ocean • Forecasting Australia Model (OFAM; 70°S–70°N domain, 1/10° resolution around Australasia)



OFAM grid with mean SST



Ocean Modeling



- Eddy-resolving dynamical downscaling in Australia region performed by Chamberlain et al. (2010):
- Two <u>ocean model runs</u> using Ocean Forecasting Australia Model (**OFAM**; 70°S–70°N domain, 1/10° resolution around Australasia)
- Forcings representative of:
 - **1990s (CTRL run)**, and
 - 2060s (A1B run)
- Control run forced by historical reanalysis
- Climate change scenario provided by CSIRO Mk3.5 GCM with an A1B emissions scenario



OFAM grid with mean SST







Ocean Modeling



- Eddy-resolving dynamical downscaling in Australia region performed by Chamberlain et al. (2010):
- Two <u>ocean model runs</u> using Ocean Forecasting Australia Model (**OFAM**; 70°S–70°N domain, 1/10° resolution around Australasia)
- Forcings representative of:
 - 1990s (CTRL run), and
 - 2060s (A1B run)
- Control run forced by historical reanalysis
- Climate change scenario provided by CSIRO Mk3.5 GCM with an A1B emissions scenario
- Models represent well general circulation and temperature distribution around Australia, including seasonality [Sun et al, 2012; Matear et al., 2013]



OFAM grid with mean SST







Mean Circulation



Model simulated surface mean dynamic topography

(indicative of surface geostrophic flow) consistent with observations







• Model simulated **mean volume transport**







• **Redistribution of flow** through the Tasman Sea



• Enhanced EAC extension and reduced flow along Tasman Front, consistent with basin-wide changes in wind stress curl





• Simple linear, wind-driven, barotropic circulation model

- If the changes in mean circulation are simply due to changes in wind-stress, then we may be able to capture them with such a model
- Sverdrup stream function $\psi(x,y)$ given by zonally integrating meridional flow according to $d\psi/dx=V$
- Wind-driven V in the interior of the ocean given by the Sverdrup balance:

$$\beta V = \frac{1}{\rho H} \nabla \times \vec{\tau}$$

- Value of ψ along island boundaries (i.e., Aus., NZ) handled by Godfrey (1989) Island Rule





• Island Rule stream function for CTRL (1990s) winds, and change for A1B (2060s) wind



circulation changes at high latitudes in the Tasman Sea (EAC extension)



Eddy Kinetic Energy





- Sea level variance (~eddy kinetic energy) consistent between model and observations
- Significant increase in eddy kinetic energy in EAC Extension region, where flow is not steady but in fact consists of a train of mesoscale eddies...



Eddy Statistics



Cyclonic (blue) and anticyclonic (red) eddies tracked using Chelton et al. (2011) sea level algorithm:



Significant increase in number of long-lived anticyclonic (warm core) eddies in EAC Extension region, and possibly an increase in eddies passing through the Tasman Leakage



The Mean SST



Model projected change to mean SST



Model projected Tasman Sea hotspot





• ...and an associated increase in **SST variance** in same region



in EAC extension region





 In 2011, a "marine heat wave" off of Western Australia was documented (Pearce and Feng, 2013; Feng et al., 2013)



 Some species experienced range extensions during the marine heat wave which persisted after the heat wave dissipated (Wernberg et al. 2013)

Extreme SSTs





- The ocean model runs **do not** fully represent the extremes
- The ocean model runs **do represent** well the overall climate
- Extremes can be represented using the "climate" alone, e.g.:
 - Griffiths et al. (2005), Ballester et al. (2010), Simolo et al. (2011), de Vries at al. (2012)





• SST Extremes modeled using a **Bayesian hierarchical model**

• For the 1990s climate:

- The CTRL run yields good estimates of the large-scale marine climate, e.g., the mean circulation and eddy variability and overall marine climate (mean SST, variance, etc)
- Model observed historical extremes using climate statistics as predictors: "extremes" = f("climate")

• For the 2060s climate:

- Use fitted model, *f("climate")*, and the projected future climate statistics to estimate future extremes
- Assumes the model is stationary in time





• <u>Bayesian hierarchical model</u>

- Data Layer
 - Assume that at each location *j* the annual maxima *y_j* are distributed according to an extreme value distribution:

$$\boldsymbol{y}_{j} \sim EVD$$
 (a_{j} , $\boldsymbol{\varphi}_{j}$)

- Climate process layer
 - Assume the model parameters over all space (*a*, φ) can be estimated by a linear regression onto a set of predictors *X*, consisting of the mean SST, SST variance, skewness, etc:

$$a = X\beta_a + \varepsilon$$
$$\phi = X\beta_{\phi} + \varepsilon$$

- The errors ($\boldsymbol{\varepsilon}$) are assumed to be normally distributed
- **Priors** are diffuse, non-informative
- Parameters are **sampled** using MCMC, Metropolis rule, Gibbs sampler

Extreme SSTs









• Projected change in extreme SSTs (50-year return levels) is due to a combination of the changes in mean and variance



Overall increase (due to change in mean SST)



Extreme SSTs



- The extremes model is probabilistic in nature (Bayesian) and so we can put confidence limits on our predictions
- This type of information is very helpful when making statements about climate change







- Can use the extremes model as a "toy model"
- Test the response of the extremes to specified changes in climate





Extreme SSTs









- Projected changes in the mean state:
 - Tasman Sea SST hotspot







- Projected changes in the mean state:
 - Tasman Sea SST hotspot
 - Redistribution of transport through Tasman Sea









- Projected changes in the mean state:
 - Tasman Sea SST hotspot
 - Redistribution of transport through Tasman Sea
- Changes to mean circulation consistent with linear, wind-driven, barotropic model









- Projected changes in the mean state:
 - Tasman Sea SST hotspot
 - Redistribution of transport through Tasman Sea
- Changes to mean circulation consistent with linear, wind-driven, barotropic model
- Changes in SST extremes are due to combination of changes in SST mean, variance, skewness, etc...











- Projected changes in the mean state:
 - Tasman Sea SST hotspot
 - Redistribution of transport through Tasman Sea
- Changes to mean circulation consistent with linear, wind-driven, barotropic model
- Changes in SST extremes are due to combination of changes in SST mean, variance, skewness, etc...
- The **Bayesian hierarchical extremes model** a general framework for estimating extremes from climate/ocean models (<u>bias correction</u>)









38

40

42

44



- Global marine climate is warming
- The SW Pacific (Tasman Sea) is a hotspot of change
- Impacts on marine ecology are already being felt
- Historical marine climate in Eastern Tasmania is poorly understood
- Large-scale models and reanalysis are <u>not designed for near-shore</u> <u>studies</u>
- Alternatives
 - Statistical downscaling
 - Dynamical downscaling







- Global marine climate is **warming**
- The SW Pacific (Tasman Sea) is a hotspot of change
- Impacts on marine ecology are already being felt
- Historical marine climate in Eastern Tasmania is poorly understood
- Large-scale models and reanalysis are <u>not designed for near-shore</u> <u>studies</u>
- Alternatives
 - Statistical downscaling
 - Dynamical downscaling

Correlation between observed (AVHRR) and reanalysis (BRAN2) SST



correlation between observed (AVHRR) and reanalysis (BRAN) SST





- Global marine climate is **warming**
- The SW Pacific (Tasman Sea) is a hotspot of change
- Impacts on marine ecology are already being felt
- Historical marine climate in Eastern Tasmania is poorly understood
- Large-scale models and reanalysis are <u>not designed for near-shore</u> <u>studies</u>
- Alternatives
 - Statistical downscaling
 - Dynamical downscaling

Correlation between observed (AVHRR) and reanalysis (BRAN2) SST



correlation between observed (AVHRR) and reanalysis (BRAN) SST
Statistical Downscaling

Oliver and Holbrook (2014; OH14) statistical technique improved estimates of shelf SST

E FOR MARINE AND







- Technique extended to the entire continental shelf for temperate Australia (>20°S)
- OH14 dataset available online: passage.phys.ocean.dal.ca/~olivere/OH14.html
- Technique has been applied to 1990s and 2060s climate projections as well





Dynamical Downscaling



- Dynamical downscaling uses numerical ocean models based on our theoretical understanding of ocean dynamics
- We modeled the eastern Tasmania continental shelf using the Sparse Hydrodynamic Ocean Code (SHOC) model [Herzfeld, 2006]
- <u>Domain</u>: South Cape to ~Eddystone
 Point and seaward out to shelf break
- <u>Bathymetry</u>: Australian Geological Survey Organisation (AGSO) 2002
- <u>Resolution</u>: ~1.9 km resolution
- 43 <u>z-levels</u> in the vertical



Herzfeld, M. (2006), An alternative coordinate system for solving finite difference ocean models, Ocean Modelling, 14 (3-4), 174-196



Dynamical Downscaling



- Dynamical downscaling uses numerical ocean models based on our theoretical understanding of ocean dynamics
- We modeled the eastern Tasmania continental shelf using the Sparse Hydrodynamic Ocean Code (SHOC) model [Herzfeld, 2006]
- <u>Domain</u>: South Cape to ~Eddystone
 Point and seaward out to shelf break
- <u>Bathymetry</u>: Australian Geological Survey Organisation (AGSO) 2002
- <u>Resolution</u>: ~1.9 km resolution
- 43 <u>z-levels</u> in the vertical



Herzfeld, M. (2006), An alternative coordinate system for solving finite difference ocean models, Ocean Modelling, 14 (3-4), 174-196



UTAS

Model forced by realistic ocean and atmosphere at boundaries

IMAS Forcing: Ocean and Atmosphere



Model forced by realistic ocean and atmosphere at boundaries

 <u>Lateral boundaries</u> were forced by velocities, temperature and salinity from **Bluelink** reanalysis and analysis fields ^{1,2}



INTAS Forcing: Ocean and Atmosphere



Model forced by realistic ocean and atmosphere at boundaries

- <u>Lateral boundaries</u> were forced by velocities, temperature and salinity from **Bluelink** reanalysis and analysis fields ^{1,2}
- <u>Surface was forcing</u> was provided from the NCEP Climate Forecast System (CFS) Reanalysis and Reforecast ^{3,4}

0.25 m/s
 3.1 m/s
 1.5 m/s
 Diffing buoy positions 15-Mar-2014
 4.5 m/s
 1.5 m/s</l



Climate Forecast System Reanalysis Annual 1979-2013 Average

¹ BRAN = Bluelink ReANalysis

- ² OceanMAPS = Bluelink Ocean Modelling, Analysis, and Prediction System
- ³ CFSR = Climate Forecast System Reanalysis

⁴ CFSv2 = Climate Forecast System version 2 (operational forecast system)

INTAS FORCING: Ocean and Atmosphere



Model forced by realistic ocean and atmosphere at boundaries

- <u>Lateral boundaries</u> were forced by velocities, temperature and salinity from **Bluelink** reanalysis and analysis fields ^{1,2}
- <u>Surface was forcing</u> was provided from the NCEP Climate Forecast System (CFS) Reanalysis and Reforecast ^{3,4}
- <u>River input</u> was reconstructed from precipitation over catchments
- <u>Coverage</u>: 1993-2013
- ¹ BRAN = Bluelink ReANalysis
- ² OceanMAPS = Bluelink Ocean Modelling, Analysis, and Prediction System
- ³ CFSR = Climate Forecast System Reanalysis

⁴ CFSv2 = Climate Forecast System version 2 (operational forecast system)





Climate Forecast System Reanalysis Annual 1979-2013 Average

INTARS FOR FORCING: Ocean and Atmosphere



Model forced by realistic ocean and atmosphere at boundaries



- <u>Lateral boun</u> velocities, te from **Bluelir** analysis fielc
- <u>Surface was</u> from the NC
 System (CFS Reforecast ^{3,}
- Four <u>hindcast simulations</u> were performed for the 1993-2013 period:
- With and without tides (T/NT)
 With and without rivers (R/NR)
- **Base run for validation** is with rivers and without tides (**R/NT**)





ClimateReanalyzer.org

Climate Change Institute | University of Maine

- <u>River input</u> v precipitation
- <u>Coverage</u>: 1993-2013
- ¹ BRAN = Bluelink ReANalysis
- ² OceanMAPS = Bluelink Ocean Modelling, Analysis, and Prediction System
- ³ CFSR = Climate Forecast System Reanalysis

⁴ CFSv2 = Climate Forecast System version 2 (operational forecast system)





In-situ time series



Maria Island time series [RED]

- Historical temperature and salinity @ surface and 5 depths
- Quasi-monthly, 1944 2008

Craig Mundy (IMAS-FAC, UTAS), near-bottom temperature gauges [BLUE]

- Near-bottom temperature in 5-20 m water depths
- Daily, 2005 present

2 Tide gauges (Hobart, Spring Bay) [BLACK]

- Sea level
- Hourly and daily, 1985 2012

<u>Remotely sensed</u>

AVHRR: daily, 4 km maps, 1980-2012



Tide Gauges



• Model captures well sea level at Hobart and Spring Bay tide gauges



Maria Island Time Series





RCTIC STUDIES

- Maria Island Time Series
- Temperature, model captures well:
 - The total variability at all depths

model

The seasonal cycle

observations

The non-seasonal variability



Maria Island Time Series





- Maria Island Time Series
- Temperature, model captures well:
 - The total variability at all depths

model

The seasonal cycle

observations -

- The non-seasonal variability
- The inter-annual variability



Maria Island Time Series





- Maria Island Time Series
- Temperature, model captures well:
 - The total variability at all depths
 - The seasonal cycle
 - The non-seasonal variability
 - The inter-annual variability
- Salinity:
 - A notable bias in mean salinity throughout the water column
 - May be related to poor representation of advection in earlier model versions

observations —— model

INTAS IN Near-bottom temperatures



Near-bottom temperature loggers

• Model captures well the total variability, seasonal cycle, and non-seasonal signal



Remotely-sensed SST



- Model captures well mean SST remotely-sensed by AVHRR
- A slight near-shore warm bias (0.5°C) in the northern 2/3 of the domain.

CTIC STUDIES

Remotely-sensed SST



- Model captures well mean SST remotely-sensed by AVHRR
- A slight near-shore warm bias (0.5°C) in the northern 2/3 of the domain.
- Remotely sensed SST variations are also well captured with correlations generally >0.85.





The Zeehan Current



- **East Australian Current (EAC)**, a quasi-steady western boundary current, separates from the coast ~33°S.
- The EAC Extension continues southward transport as far as Tasmania, but as an unsteady, eddy-rich "current"
- The **Zeehan Current**, part of a current system extending all the way to WA, runs southward and eastward along the west and south coasts of Tasmania [Ridgway and Condie, 2004]



http://www.marine.csiro.au/~lband/yacht_races/yyzeecur.html



 Along the southeast coast of Tasmania, the EAC Extension is dominant in summer and the Zeehan Current is dominant in winter



Surface Climatology







Seasonal Climatology



34.5





Seasonal Climatology







Roles of tides and rivers



- Role of the rivers (R/NT NR/NT):
 - Reduced salinity in Derwent and Huon estuaries
 - Presence of Derwent cools while the Huon warms (not shown)

- <u>Tidal interactions</u> (*R/T R/NT*):
 - Tide-River interactions can be significant in and around river estuaries, Note: <u>Frederick Henry</u> <u>Bay and Norfolk Bay</u>







- Composite averages show a strong signature of Southern Annular Mode (SAM) on marine climate
- +SAM leads to increased salinity in the Derwent and Huon estuaries while -SAM leads to decreased salinity







- Composite averages show a strong signature of Southern Annular Mode (SAM) on precipitation over the catchment
- +SAM leads to decreased rainfall while -SAM leads to increased rainfall over SW Tasmania







- **ETAS** model compares well against observed coastal sea level, historical Maria Island time series, near-bottom temperature at a number of sites, and remotely-sensed SST across the shelf
- Model captures seasonal alternation between dominance of Zeehan Current in Winter (JJA) and EAC Extension in Summer (DJF)
- <u>Roles of rivers, tides, and climate modes</u>:
 - **Rivers** freshen and cool estuary waters
 - **Tides** can interact with the rivers in a complex way
 - Climate mode (SAM) plays a role in modulating near-shore marine climate through precipitation and thus river inputs
- **Future work**: relative role of surface and boundary forcing, interaction between off-shore eddies and the shelf, influence of ENSO and Tasman Sea blocking, modulation of shelf circulation by climate modes



Acknowledgements



ARC CENTRE OF EXCELLENCE FOR

CLIMATE SYSTEM SCIENCE

- Neil Holbrook (IMAS) post-doc supervisor
- Simon Wotherspoon (IMAS) statistics mathemagician
- Matthew Chamberlain, Richard Matear (CSIRO) OFAM modelling
- CSIRO EMS Team: Mike Herzfeld, John Andrewartha, Mark Baird, Farhan Rizwi
- SSF team: Martin Marzloff, Craig Johnson, and Neville Barrett (IMAS)
- Craig Mundy (IMAS) Tassie temperature logger data
- Mauro Vargas Hernandez, Andre Belo do Couto, Bo Qiu, Keith Thompson, Jessica Benthuysen, Max Nikurashin, Andrew Kiss, Terry O'Kane, etc... for helpful discussions
- The Super Science Fellowship and the Centre of Excellence for Climate System Science (Australia Research Council), IMAS and UTAS for support, financial and otherwise



Australian Research Council



Acknowledgements



ARC CENTRE OF EXCELLENCE FOR

CLIMATE SYSTEM SCIENCE

- Neil Holbrook (IMAS) post-doc supervisor
- Simon Wotherspoon (IMAS) statistics mathemagician
- Matthew Chamberlain, Richard Matear (CSIRO) OFAM modelling
- CSIRO EMS Team: Mike Herzfeld, John Andrewartha, Mark Baird, Farhan Rizwi
- SSF team: Martin Marzloff, Craig Johnson, and Neville Barrett (IMAS)
- Craig Mundy (IMAS) Tassie temperature logger data
- Mauro Vargas Hernandez, Andre Belo do Couto, Bo Qiu, Keith Thompson, Jessica Benthuysen, Max Nikurashin, Andrew Kiss, Terry O'Kane, etc... for helpful discussions
- The Super Science Fellowship and the Centre of Excellence for Climate System Science (Australia Research Council), IMAS and UTAS for support, financial and otherwise



Australian Research Council

Publications

Oliver, E. C. J. and N. J. Holbrook (2014), Extending our understanding of South Pacific gyre 'spin-up': Modeling the East Australian Current in a future climate, *Journal of Geophysical Research*, 119, 2788-2805

Oliver, E. C. J., S. J. Wotherspoon and N. J. Holbrook (2014), Estimating extremes from global ocean and climate models: A Bayesian hierarchical model approach, *Progress in Oceanography*, 122, 77-91

Oliver, E. C. J., S. J. Wotherspoon, M. A. Chamberlain and N. J. Holbrook (2014), Projected Tasman Sea extremes in sea surface temperature through the 21st century, *Journal of Climate*, 27(5), 1980-1998

Oliver, E. C. J. and N. J. Holbrook (2014), A statistical method for improving continental shelf and nearshore marine climate prediction, *Journal of Atmospheric and Oceanic Technology*, 31, 215-232





Extra Slides...





- The third moment of sea level (sea level skewness) can be used to map the mean path of a meandering jet, such as western boundary currents [Thompson and Demirov, 2006]
- The intersection of the mean jet path with the shelf break is used as a rough indicator of the mean EAC separation point



Approx 90 km southward shift in EAC separation

Changes at Depth





- Changes present throughout the water column: a general warming and freshening of the Tasman Sea
- Deepening of EAC anticyclonic recirculation near separation point?







 Surface changes along the shelf break indicate a consistent warming of ~2°C and freshening only north of Bass Strait and increasing with latitude



Model Stationarity

Fundamental relationship

We posit that there exists a relationship between the extremes and climate parameters X:

"extremes" =
$$f(\mathbf{X})$$

This relationship expresses fundamental aspects of the climate system which do not change with time.

Role of β s and τ s

Effectively, we have performed a linear approximation to $f(\mathbf{X})$:

 $f(\mathbf{X}) = \mathbf{X}\boldsymbol{\beta} + O(\mathbf{X}^2)$

Therefore, the β s (and τ s) are stationary since $f(\mathbf{X})$ is stationary

DQ P



River Input



- We also require river input (<u>flow rate and water temperature</u>) for the two major rivers in SE Tasmania: Derwent River and Huon River
- We have <u>observed records</u> of flow (m³/s) and water temp for both rivers, but records very short and very recent (Nov/2009 -late/2013; shorter for temp) and we require these quantities over the entire 1993-2013 period
- Therefore, we <u>modeled river flow</u> (F) using precipitation (P) over the river catchments (from CFSR/CFSv2) as a predictor in a multiple lag-regression model:

$$\log(F_t) = \alpha + \sum_{l=0}^{L} \beta_l \log(P_{t-l})$$

And a similar model (without log-transforms) to estimate river temperature from local air temperature

- A two-fold cross-validation was performed to determine which value of L provided the best fit
- Given a satisfactory fit, we used historical precipitation and air temperature from CFSR/CFSv2 to <u>reconstruct river flow and temperature</u> over the entire 1993-2013 period



River Input



- River input (<u>flow rate and water temperature</u>) required for Derwent River and Huon River
- River inputs predicted from precipitation and air temperature using a lagregression model and then reconstructed over 1993-2013







- <u>Spin-up</u>: model was spun up for 5 years using normal year forcing*, initialized from Bluelink on 1/1/1993
- <u>Historical hindcast</u>: model was then forced by realistic forcing over the 1993-2013 period
- <u>Four runs</u> were performed for all combinations
 - with and without tidal forcing (T and NT), and
 - with and without river inputs (R and NR)
 - The base run for validation was with river input and no-tides (R/NT)

* NYF: climatological seasonal cycle, subseasonal variability from 1995, mean of 1993; following *Large and Yeager (2004)*; rivers are climatological only





- Composites of surface salinity modes of climate variability show a strong signature of ENSO and SAM on marine climate
- +SAM and La Nina lead to increased SSS in the Derwent and Huon estuaries while -SAM and El Nino lead to decreased SSS
- Signal disappears in NR runs: must be **related to river input**
- SAM and ENSO are related to a +/- 6–8 m³/s modulation of river flow, which is up to 10% of the mean flow


Modes of Variability



RCTIC STUDIES

